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Fluor-rewitzerite, [(H₂O)K]Mn₂(Al₂Ti)(PO₄)₄(OF)(H₂O)₁₀ · 4H₂O, a new paulkerrite-group mineral, from the Hagendorf-Süd pegmatite, Oberpfalz, Bavaria, Germany

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Abstract. Fluor-rewitzerite, $[(H_2O)K]Mn_2(Al_2Ti)(PO_4)_4(OF)(H_2O)_{10} \cdot 4H_2O$, is a new monoclinic member of the paulkerrite group, from the Hagendorf-Süd pegmatite, Oberpfalz (Upper Palatinate in English), Bavaria, Germany. It occurs on the walls of vugs in corroded zwieselite, in association with Zn- and Al-bearing earlshannonite, fluorapatite, jahnsite-(CaMnMn) and Al-rich strunzite. Fluor-rewitzerite forms clusters of colourless stubby prisms up to 0.1 mm long that are flattened on {010}; elongated along [100]; and show the forms {100}, {010}, {001}, {111} and {111}. Twinning occurs by 2-fold rotation about *c*. The measured density is 2.42(2) g cm⁻³. Optically, fluor-rewitzerite crystals are biaxial (+), with $\alpha = 1.569(3)$, $\beta = 1.582(3)$, $\gamma = 1.602(3)$ (white light) and $2V(\text{meas}) = 78(1)^\circ$. The empirical formula from electron microprobe analyses and structure refinement is ${}^{A1}[(H_2O)_{0.85}K_{0.15}]_{\Sigma 1.00}{}^{A2}(K_{1.00}){}^{M1}(Mn^{2+}_{1.50}Mg_{0.09}Fe^{2+}_{0.41})_{\Sigma 2.00}{}^{M2+M3}(Al_{1.70}Ti^{4+}_{0.89}Fe^{3+}_{0.42})_{\Sigma 3.01}(PO_4)_{3.99}{}^{X}(O_{1.09}F_{0.92})_{\Sigma 2.01}(H_2O)_{10} \cdot 4.12H_2O$. Fluor-rewitzerite has monoclinic symmetry with space group $P2_1/c$ and unit-cell parameters a = 10.407(1) Å, b = 20.514(2) Å, c = 12.193(1) Å, $\beta = 90.49(2)^\circ$, V = 2603.0(4) Å³ and Z = 4. The crystal structure was refined using synchrotron single-crystal data to $R_{obs} = 0.058$ for 6186 reflections with $I > 3\sigma(I)$. Fluor-rewitzerite is the fluoride analogue of rewitzerite, with F dominant over OH at the X sites of the general formula $A1A2M1_2M2_2M3(PO_4)_4X_2(H_2O)_{10} \cdot 4H_2O$.

1 Introduction

Rewitzerite, $[K(H_2O)]Mn_2(Al_2Ti)(PO_4)_4[O(OH)](H_2O)_{10} \cdot 4H_2O$, was recently approved as a new monoclinic mineral of the paulkerrite group (Grey et al., 2023a). The ideal formula for rewitzerite conforms to the general formula for the monoclinic group members, $A1A2M1_2M2_2M3(PO_4)_4X_2(H_2O)_{10} \cdot 4H_2O$, where A = K, H_2O , \Box (denoting a vacancy); $M1 = Mn^{2+}$, Mg, Fe²⁺, Zn (rarely Fe³⁺); M2 and $M3 = Fe^{3+}$, Al, Ti⁴⁺; and X = O,

OH, F. A high degree of mixing of Fe³⁺, Al and Ti⁴⁺ occurs at the *M*2 and *M*3 sites, and charge balance is maintained for different M^{3+}/M^{4+} ratios at *M*2 and *M*3 by variations in the X^{-}/X^{2-} ratios at the *X* sites. For end-member rewitzerite, with *M*2₂*M*3 = Al₂Ti, co-dominant OH⁻ and O²⁻ are at the *X* sites.

In ongoing studies of potential new secondary phosphate minerals in the Erich Keck collection at the Bavarian State Collection for Mineralogy, Munich, using scanning electron microscopy coupled with energy-dispersive X-ray analysis, we identified a specimen as having an analysis very similar to that for rewitzerite but with more than twice the F content. More detailed studies, including electron microprobe analyses and single-crystal X-ray diffraction, confirmed the mineral as the fluoride analogue of rewitzerite with F dominant over OH at the X sites in the crystal structure. The mineral and its name have been approved by the International Mineralogical Association (IMA) Commission on New Minerals, Nomenclature and Classification (CNMNC), proposal IMA-2023-115.

The holotype specimen is housed in the mineralogical collections of the Bavarian State Mineral Collection, Munich, registration number MSM 80824. A cotype specimen used for the optical properties, powder X-ray diffraction (XRD) and Raman spectrum is located at the Natural History Museum of Los Angeles County, catalogue number 76307.

2 Occurrence and associated minerals

Fluor-rewitzerite was found in a recent study of specimens of altered zwieselite that were collected by Erich Keck in the mid-1970s from the 67 m level of the Hagendorf-Süd pegmatite mine quarry in the Oberpfalz (Upper Palatinate in English), northeast Bavaria (49°39'1" N, 12°27'35" E). The holotype specimen consists of altered light-brown zwieselite with many vugs up to 5 mm in diameter. The vugs are lined with tiny (< 0.1 mm) colourless crystals of fluor-rewitzerite sitting directly on the corroded zwieselite. Associated minerals are fluorapatite, white needles of Al-rich strunzite, tabular yellow crystals of Na-bearing jahnsite-(CaMnMn) and vellow microspheres of Zn- and Al-bearing earlshannonite. A massive black mineral filling some vugs has been identified by powder XRD as a Zn-bearing cryptomelane. Directly sitting on the crystals of fluor-rewitzerite are tiny white balls of iangreyite as the youngest member of the association. Directly embedded in the zwieselite and sometimes protruding into the corrosion vugs - and thus older than the fluor-rewitzerite - are massive sphalerite, tabular crystals of columbite and small grains of pyrite.

3 Physical and optical properties

Fluor-rewitzerite forms clusters of colourless stubby prisms up to 0.1 mm long (Figs. 1 and 2). The crystals are flattened on {010}; elongated along [100]; and show the forms {100}, {010}, {001}, {111} and {111} (Fig. 2). Twinning occurs by 2-fold rotation about *c*. The density, measured by the flotation method in mixtures of methylene iodide and toluene, is $2.42(2) \text{ g cm}^{-3}$, compared with a value of 2.418 g cm^{-3} calculated for the empirical formula and single-crystal unit-cell parameters.

Optically, fluor-rewitzerite crystals are biaxial (+), with $\alpha = 1.569(3)$, $\beta = 1.582(3)$ and $\gamma = 1.6021(3)$ (measured in white light). The measured 2V from extinction data analysed



Figure 1. Fluor-rewitzerite crystals in a vug in altered zwieselite. The field of view (FOV) is 0.60 mm. Photo by Christian Rewitzer.

with EXCALIBR (Gunter et al., 2004) is $78(1)^{\circ}$, and the calculated 2*V* is 78.6°. Dispersion was not observed. The optical orientation is X = b, Y = c and Z = a. The Gladstone–Dale compatibility index (Mandarino, 1981) is 0.035 (excellent) based on the empirical formula and the measured density.

4 Chemical composition

Crystals of fluor-rewitzerite were analysed using wavelength-dispersive electron microprobe (EMP) spectrometry on a JEOL JXA-8530F Hyperprobe operated at an accelerating voltage of 15 kV and a beam current of 2.0 nA. The beam was defocused to between 8 and 10 μ m. Both specimen and standards were coated with a 25 Å thick film of iridium for the analyses. Highly hydrated paulkerrite-group minerals present problems for analysis because of dehydration in the high vacuum of the microprobe, resulting in severe cracking and higher analytical totals (Sejkora et al., 2006). To minimise this effect, a cold stage cooled to liquid nitrogen temperature was employed in the microprobe and the specimen was precooled under dry nitrogen prior to introduction to the microprobe vacuum.

Analytical results (average of 13 analyses on four crystals) are given in Table 1. There was insufficient material for direct determination of H₂O, so it was based upon the crystal structure and stoichiometry (15 H₂O per 4 P). The calculated water content was iterated in the matrix correction procedure. The EMP results show a negative correlation between Al₂O₃ and Fe₂O₃ ($R^2 = 0.66$) and a weak positive correlation be-



Figure 2. (a) Fluor-rewitzerite crystal, FOV 0.10 mm (photo Christian Rewitzer). (b) Crystal drawing of fluor-rewitzerite crystal.

tween Al₂O₃ and F ($R^2 = 0.40$). No significant correlations were found for K and Ti with other elements. The analysis results for rewitzerite (Grey et al., 2023a) are included in Table 1 for comparison. Fluor-rewitzerite has 50 % more K₂O than, 80 % more MnO than and more than twice the F content of rewitzerite but has lower TiO₂ and MgO contents.

The atoms per formula unit (apfu), normalised to 9(M1 + M2 + M3 + P) apfu, are as follows:

$$K_{1.15}Mn_{1.50}Mg_{0.09}Fe^{2+}_{0.41}Fe^{3+}_{0.42}Al_{1.70}Ti^{4+}_{0.89}P_{3.99}F_{0.92}$$

O_{32.02}H_{29.94}.

Expressing the apfu in structural form (general formula given in the Introduction) and grouping the M2 and M3 sites using the merged site procedure described by Grey et al. (2023b) gives the following empirical formula:

$${}^{A1}[(H_2O)_{0.85}K_{0.15}]_{\Sigma 1.00} {}^{A2}(K_{1.00})$$

$${}^{M1}(Mn^{2+}{}_{1.50}Mg_{0.09}Fe^{2+}{}_{0.41})_{\Sigma 2.00}$$

$${}^{M2+M3}(Al_{1.70}Ti^{4+}{}_{0.89}Fe^{3+}{}_{0.42})_{\Sigma 3.01}(PO_4)_{3.99}$$

$${}^{X}(O_{1.09}F_{0.92})_{\Sigma 2.01}(H_2O)_{10} \cdot 4.12H_2O.$$

The simplified formula is

The ideal formula is $[(H_2O)K]Mn_2(Al_2Ti)(PO_4)_4(OF)(H_2O)_{10}$. 4H₂O, which requires K₂O 5.03, MnO 15.16, Al₂O₃ 10.90, P₂O₅ 30.34, TiO₂ 8.54, F 2.03, H₂O 28.85, $-O \equiv F - 0.85$ and total 100.00 wt %.

Note that the merged $(M2_2M3)$ sites approach is illustrated graphically in Fig. 4, showing a ternary Al₃–Fe₃³⁺–Ti₃ diagram with the possible end-member compositions designated (Ti₃, Al₂Ti, Ti₂Al, etc.) and the empirical compositions

for published paulkerrite-group minerals shown by crosses. The empirical $(M2_2M3)$ composition for fluor-rewitzerite is shown in Fig. 4 to be located in the end-member (Al₂Ti) composition field. Combining this with the dominant species at the *A*, *M*1 and *X* sites leads to the end-member (ideal) formula reported above.

5 Raman spectroscopy

Raman spectroscopy was conducted on a Horiba XploRA PLUS spectrometer using a 532 nm diode laser, 100 µm slit, 1800 grooves mm⁻¹ diffraction grating and a $100 \times (0.9 \text{ nu})$ merical aperture) objective. The spectrum, shown in Fig. 3, is almost identical to that for rewitzerite and the same assignments apply as given for rewitzerite (Grey et al., 2023a). The O-H stretch region has a broad peak that can be assigned to H-bonded water, with a maximum at 3355 cm^{-1} and a shoulder at 3115 cm^{-1} . Hydroxyl ion stretching is evident by a shoulder at 3570 cm^{-1} . The H–O–H bending mode region for water has a peak at $1650 \,\mathrm{cm}^{-1}$. Three strong peaks at 955, 970 and 1015 cm^{-1} in the P–O stretching region can be assigned to symmetric stretching modes, while weaker peaks at 1095 and 1130 cm⁻¹ correspond to antisymmetric P-O stretching modes. Bending mode vibrations of the $(PO_4)^{3-}$ groups are located at 600 cm^{-1} and at 425, 445 and $480 \,\mathrm{cm}^{-1}$. Peaks at lower wavenumbers are related to lattice vibrations. A strong peak at 835 cm^{-1} with a shoulder at 780 cm⁻¹ can be assigned to Ti–O stretch vibrations for short Ti-O bonds that occur in linear trimers of cornerconnected octahedra M2-M3-M2 in the structure, by analogy with published Raman spectra for titanates containing short Ti-O distances (Tu et al., 1996; Bamberger et al., 1990; Silva et al., 2018).

	Fluor-rewitzerite, average of 13 analyses			Rewitzerite (Grey et al., 2023a)	
Constituent	Mean	Range	SD	Mean (SD)	Standard
K ₂ O	5.78	5.40-6.22	0.20	3.93 (0.61)	Adularia
MnO	11.35	9.64-12.44	0.91	6.33 (0.78)	MnSiO ₃
MgO	0.42	0.23-0.52	0.08	2.80 (0.52)	Spinel
Al_2O_3	9.27	6.20-11.14	1.35	8.37 (1.87)	Berlinite
Fe ₂ O ₃ (total)	7.01	4.04-9.81	1.99	7.44 (1.59)	Hematite
Fe ₂ O ₃ ^a	3.54				
FeO ^a	3.11				
TiO ₂	7.64	6.93-8.94	0.60	9.18 (0.77)	Rutile
P_2O_5	30.28	28.42-31.61	1.02	30.90 (2.42)	Berlinite
F	1.88	1.18-2.42	0.40	0.87 (0.31)	Fluorite
H ₂ O _{calc} ^b	28.80			30.39	
$-0 \equiv F$	-0.79			-0.37	
Total	101.28			99.84	

 Table 1. Analytical data (wt %) for fluor-rewitzerite.

^a Based on Fe²⁺ at M1 and Fe³⁺ at the M2 and M3 sites. ^b Based on stoichiometry: 15 H₂O per 4 P.



Figure 3. Raman spectrum of (a) fluor-rewitzerite and (b) rewitzerite (Grey et al., 2023a).

6 Crystallography

6.1 X-ray powder diffraction

X-ray powder diffraction data were recorded using a Rigaku R-Axis Rapid II curved-imaging-plate microdiffractometer with monochromatised MoK α radiation. A Gandolfilike motion on the ϕ and ω axes was used to randomise the sample. Observed *d* values and intensities were derived by profile fitting using JADE Pro software (Materials Data, Inc.). Data are given in Table 2. Refined monoclinic unit-cell parameters (space group $P2_1/c(\text{no. 14})$) are a = 10.425(18) Å, b = 20.591(18) Å, c = 12.285(18) Å, $\beta = 90.47(10)^\circ$, V = 2637(6) Å³ and Z = 4.



Figure 4. Ternary diagram for $(M2)_2M3$ site Al–Ti–Fe³⁺ compositions, showing end-member compositions (Al₂Ti, Ti₂Al, etc.) and location of the empirical composition for fluor-rewitzerite, For comparison the published empirical compositions are shown for the paulkerrite-group minerals benyacarite (Demartin et al., 1997), hochleitnerite (Grey et al., 2023d), hydroxylbenyacarite (Hochleitner et al., 2024), mantienneite (Fransolet et al., 1984), paulkerrite (Peacor et al., 1984), pleysteinite (Grey et al., 2023c), rewitzerite (Grey et al., 2023a) and macraeite (Grey et al., 2024). Red crosses correspond to minerals with Mn at *M*1, and blue crosses correspond to minerals with Mg at *M*1.

6.2 Synchrotron single-crystal diffraction

A crystal measuring $0.040 \times 0.050 \times 0.070 \text{ mm}^3$ was used for a single-crystal diffraction data collection at the Australian Synchrotron microfocus beamline MX2 (Aragao et al., 2018). Intensity data were collected using a Dectris Eiger 16M detector and monochromatic radiation with a wave-

Table 2. Powder X-ray diffraction data (d in Å) for fluor-rewitzerite ($I_{calc} > 1.5$).

Table 2. Continued.

I _{obs}	d _{obs}	d_{calc}	<i>I</i> _{calc}	hkl
		10.481	2	011
46	10.256	10.257	40	020
		9.281	2	110
62	7.414	7.359	57	111
		6.282	2	-121
100	6.149	6.096	100	002
41	5.190	5.203	24	200
		5.166	3	131
		5.129	11	040
17	4.659	4.640	12	220
32	3.955	3.974	5	-202
		3.913	26	231
64	3.704	3.733	4	-113
		3.705	21	-222
		3.680	19	222
		3.653	8	240
7	3.407	3.419	6	060
		3.319	2	-133
74	3.119	3.111	50	251
37	3.036	3.048	20	004
		3.025	6	-302
		3.005	6	302
24	2.962	2.982	11	062
_		2.932	9	-104
80	2.866	2.919	11	104
		2.902	2	-322
		2.883	2	322
		2.857	32	260
14	2.812	2.819	5	-124
		2.808	5	124
		2.786	12	-153
45	2.603	2.639	3	-204
		2.621	4	204
		2.606	9	-342
		2.593	16	342
		2.583	3	262
		2.556	2	-224

Iobs	$d_{\rm obs}$	d_{calc}	<i>I</i> _{calc}	hkl
34	2.519	2.545	5	-144
		2.529	10	-411
		2.521	4	411
		2.501	12	-271
		2.470	2	-333
		2.453	3	333
		2.364	2	082
20	2.323	2.337	3	-422
		2.324	2	422
		2.303	9	182
		2.275	5	064
17	2.176	2.174	6	-442
		2.163	7	442
		2.155	2	-282
7	2.095	2.089	2	-264
		2.080	2	264
		2.070	3	460
10	2.053	2.057	4	291
		2.032	8	006
29	1.972	1.987	4	-404
		1.971	3	404
		1.964	9	-462
		1.957	9	462
		1.951	2	382
		1.939	2	-5 2 2
13	1.915	1.923	2	-471
		1.910	3	1 10 2
		1.889	6	046
9	1.866	1.866	4	-226
		1.857	5	226
		1.851	2	-513
15	1.835	1.839	7	513
		1.833	2	284
		1.823	7	-2 10 2
9	1.771	1.767	2	-275
		1.760	5	-306
		1.747	5	306

Table 2. Continued.

Iobs	d _{obs}	d _{calc}	<i>I</i> _{calc}	hkl
15	1.716	1.718	2	-464
		1.710	12	0 12 0
		1.697	2	491
		1.679	2	571
28	1.653	1.664	2	-346
		1.660	3	-266
		1.653	6	266
		1.643	13	640
		1.624	3	2 12 0
18	1.600	1.611	5	4 10 0
		1.593	2	086
		1.588	6	-426
		1.583	5	642
10	1.561	1.576	3	426
		1.559	3	-4 10 2
		1.556	3	4 10 2
		1.547	3	660
6	1.511	1.525	2	-475
		1.503	4	4 11 1
		1.491	2	0 12 4
17	1.462	1.468	2	-1 11 5
		1.461	4	048
14	1.440	1.451	2	-644
		1.444	7	0 10 6
		1.437	3	680
		1.422	2	-4 11 3
		1.410	2	2 14 0
		1.393	2	-2 10 6
		1.389	2	2 10 6
		1.372	2	-3 12 4
6	1.320	1.315	3	-2 15 1
		1.311	2	408
10	1.299	1.296	2	-6 10 2
		1.292	4	-811
		1.290	2	811
		1.282	2	0 16 0

length of 0.7109 Å. The crystal was maintained at 100 K in an open-flow nitrogen cryostream during data collections. The diffraction data were collected using a single 36 s sweep of 360° rotation around ϕ . The resulting dataset consists of 3600 individual images with an approximate ϕ angle of each image being 0.1°. The raw intensity dataset was processed using XDS software to produce data files that were analysed using SHELXT (Sheldrick, 2015) and JANA2006 (Petříček et al., 2014). Refined unit-cell parameters and other data collection details are given in Table 3.

A structural model for fluor-rewitzerite was obtained in space group $P2_1/c$ using SHELXT (Sheldrick, 2015). The model had the same structure as for rewitzerite (Grey et al., 2023a), and so the rewitzerite coordinate file was used to initiate the refinement to ensure the same atom labelling. Twinning was implemented with 2-fold rotation about c. Consistent with the model for rewitzerite, Mn plus Mg was incorporated at the M1 sites, Al and Ti at the M2 and M3 sites, and K and O (for H_2O) at the A sites, and their occupancies were refined to obtain the site scattering. Refinement of the K/O ratio at the A sites gave K occupancies lower than the K content analysed by electron microprobe (EMP). A likely explanation for this discrepancy is that the A sites also contain vacancies, as reported for other paulkerrite-group minerals (Hochleitner et al., 2024; Rewitzer et al., 2024). Adjusting the vacancy content gave a reasonable match to the EMPderived K content for 0.1 vacancies per A site.

Difference Fourier maps were used to locate H atoms. A total of 25 H atoms were unambiguously located of a possible 30 H atoms from 15 independent H₂O groups. The H atoms were refined with soft restraints (O–H = 0.85(1) Å and H–O–H = $109.47(5)^{\circ}$).

Refinement with anisotropic displacement parameters for all non-H atoms and an overall isotropic parameter for Hs in JANA2006 converged at $R_{obs} = 0.058$ for 6186 reflections with $I > 3\sigma_I$. Details of the data collection and refinement are given in Table 3. The refined coordinates, equivalent isotropic displacement parameters and bond valence sum (*BVS*) values (Gagné and Hawthorne, 2015) from the refinement are reported in Table 4. Refined site scattering values are in Table 5. Anisotropic displacement parameters for the metal atoms are given in Table 6, and H bonds are reported in Table 7. Selected interatomic distances are reported in Table 8.

7 Discussion

The crystal structure for fluor-rewitzerite can be described in terms of a structural unit plus an interstitial complex, using the terminology of Hawthorne (1992) and Hawthorne and Schindler (2008). The structural unit is a 3D anionic framework of corner-connected octahedra and tetrahedra, with formula $[Mn_2Al_2Ti(PO_4)_4(OF)(H_2O)_{10}]^{1-}$, while the interstitial complex is $[K(H_2O)(H_2O)_4]^{1+}$, where the first H₂O corTable 3. Crystal data and structure refinement for fluor-rewitzerite.

$[(H_2O)K]Mn_2(Al_2Ti)(PO_4)_4(OF)(H_2O)_{10}\cdot 4H_2O$
935.7
Monoclinic, $P2_1/c$ (#14)
a = 10.407(1) Å
b = 20.514(2) Å
c = 12.193(1) Å
$\beta = 90.49(2)^{\circ}$
2603.0(4) Å ³ , 4
Synchrotron microfocus beamline
0.7109 Å, 100 K
$0.040 \times 0.050 \times 0.070 \mathrm{mm}^3$
Multiscan, T_{\min} 0.42, T_{\max} 0.75
2-fold rotation about <i>c</i>
Twin volumes 0.800(3), 0.200(3)
1.94 to 32.05°
$-13 \le h \le 13, -28 \le k \le 28, -16 \le l \le 16$
47 535
6931
6186
Full-matrix least-squares method on F
6931/89/477
$R_{\rm obs} = 0.058, \mathrm{w}R_{\rm obs} = 0.076$
$R_{\rm obs} = 0.062, \mathrm{w}R_{\rm obs} = 0.077$
0.99 and $-1.25 \text{ e} \text{ Å}^2$

responds to transformer-type O15a, coordinated to K, and the four other H₂O groups are involved only in H bonding. These comprise O14a, O14b, O15b and A1 in Tables 4 and 8. The structural unit is based on (001) heteropolyhedral layers at z = 1/4 and 3/4, as shown in Fig. 5. The layers are built from [100] kröhnkite-type chains (Hawthorne, 1985) of four-member rings of corner-connected PO₄ tetrahedra and $M2O_4X(H_2O)$ octahedra. Each PO₄ tetrahedron also shares a corner with $M1O_2(H_2O)_4$ octahedra along [010]. The corner-shared linkages form eight-member rings of alternating octahedra and tetrahedra. Layers as shown in Fig. 5 are interconnected into an open 3D framework by cornersharing of the $M2O_4X(H_2O)$ octahedra with $M3O_4X_2$ octahedra located at z = 0 and 1/2.

The interstitial complex is shown in Fig. 6. It corresponds to a (001) section through the structure centred at z=0or z=1/2 with a thickness of ~2.7 Å. In addition to the interstitial constituents and the *M*3-centred octahedra, the section contains H₂O groups that are coordinated to *M*1 cations (O10 and O12) and to *M*2 cations (O13). In-section H bonding is shown by dashed red lines on the left-hand side of Fig. 6, and in-section bonds to *A*1 and *A*2 are shown by the blue lines. Linear chains of regularly spaced water molecules, with a mean separation of ~2.9 Å, are aligned along [110] as shown by the arrows in Fig. 6. The H bonding contributes to meeting the local bond valence requirements of undersaturated O anions shared between *M*-centred octahedra and PO₄ tetrahedra. These are O1 to O8 in Table 4.



Figure 5. The (001) section through the fluor-rewitzerite crystal structure at z = 1/4.

Only O4a and O4b, shared between the *M*2-centred octahedra and PO₄, are not acceptors for H bonds as their valence requirements are essentially met by *M*2 and P (see *BVS* in Table 4). In addition to the H bonds to O1 to O8, there are extensive H bonds between water molecules in the interstitial complex. The strongest is from O13 to O15, with an O13…O15 distance of ~ 2.7 Å, shown in Fig. 6 and in Table 7. The O13a and O13b water molecules also form H bonds with three acceptor anions. In these cases, the H points towards a face

Atom	x	У	Z	Ueq	BVS
Mla	0.49393(4)	0.746848(17)	0.24525(3)	0.01738(16)	2.14
<i>M</i> 1b	0.99657(4)	-0.246861(17)	-0.24225(3)	0.01838(16)	2.13
M2a	0.65985(8)	0.50245(3)	0.74352(7)	0.0192(3)	3.53
M2b	1.16020(7)	-0.00241(2)	-0.74189(6)	0.0195(3)	3.49
M3a	0.5	0.5	0.5	0.0189(4)	3.43
<i>M</i> 3b	1	0	-0.5	0.0177(4)	3.46
P2a	0.58641(7)	0.58909(3)	0.29617(6)	0.01632(19)	4.98
P2b	1.08598(7)	-0.08934(3)	-0.29820(6)	0.01636(19)	4.98
P1a	0.90775(7)	0.59363(3)	0.80206(6)	0.01613(19)	4.99
P1b	1.40876(7)	-0.09344(3)	-0.80349(6)	0.01669(19)	5.00
A1	0.7162(2)	0.85199(9)	0.06147(17)	0.0314(7)	0.11
A2	1.22167(8)	-0.35426(3)	-0.05779(6)	0.0219(2)	0.85
X1	0.6388(2)	0.50205(7)	0.59864(17)	0.0143(5)	1.38
X2	1.1409(2)	-0.00368(7)	-0.59740(17)	0.0142(5)	1.34
O1a	0.9045(2)	0.66779(9)	0.80680(17)	0.0209(6)	1.74
Olb	1.4032(2)	-0.16787(9)	-0.80667(17)	0.0215(6)	1.72
O2a	1.0278(2)	0.57085(9)	0.73901(16)	0.0183(5)	1.77
O2h	1.5279(2)	-0.07012(9)	-0.73874(16)	0.0176(5)	1.75
03a	0.90910(20)	0.56544(9)	0.91916(15)	0.0175(5)	1.82
O3h	14116(2)	-0.06628(9)	-0.92072(16)	0.0170(5)	1.82
04a	0.7852(2)	0.00020(9) 0.57029(9)	0.7298(16)	0.0180(5) 0.0184(5)	1.87
O4h	1.2858(2)	-0.06974(9)	-0.74594(16)	0.0187(5)	1.07
010	0.5946(2)	0.66330(9)	0.71371(10) 0.29278(18)	0.0109(5) 0.0214(6)	1.71
05u 05b	1.0936(2)	-0.16354(9)	-0.29276(10)	0.0211(0)	1.71
050	0.4653(2)	-0.1033+(9) 0.56582(9)	-0.29340(17) 0.23530(16)	0.0200(5)	1.75
O6h	0.4055(2)	-0.06573(9)	-0.23735(16)	0.0173(5)	1.85
079	0.5000(2) 0.5855(2)	0.56526(9)	0.23733(10) 0.41524(16)	0.0180(5)	1.05
07a 07b	1.0828(2)	0.30520(9)	0.41924(10) 0.41810(15)	0.0130(5)	1.04
0%	0.7081(2)	-0.00047(9) 0.56246(9)	-0.41019(15) 0.23852(16)	0.0177(5)	1.99
OSh	1.2079(2)	-0.06239(9)	-0.23032(10)	0.0193(5)	1.83
000	0.3451(2)	-0.00239(9) 0.68455(10)	-0.24172(10) 0.17428(18)	0.0192(5)	0.30
09a 09b	0.3451(2) 0.8464(2)	-0.18755(10)	-0.17420(10)	0.0223(0) 0.0283(6)	0.39
010	0.0404(2) 0.5778(2)	-0.13755(10)	-0.1744(2) 0.07856(18)	0.0205(0)	0.37
010a	1.0857(2)	0.74233(10) 0.24137(10)	0.07850(18)	0.0225(0)	0.54
0100	0.6428(2)	-0.24137(10) 0.80796(10)	-0.07409(10) 0.31705(18)	0.0225(0)	0.40
011h	1.1473(2)	-0.30809(10)	-0.3057(2)	0.0220(0)	0.33
0120	0.3070(2)	-0.30009(10) 0.75123(10)	-0.5057(2)	0.0275(0)	0.40
012a 012b	0.3979(2) 0.0027(2)	0.75123(10) 0.25300(10)	0.40394(18) 0.40046(18)	0.0230(0)	0.44
0120	0.9027(2)	-0.23309(10) 0.50400(8)	-0.40940(18)	0.0230(0)	0.31
013h	1.1632(3)	-0.00056(8)	-0.9113(2)	0.0212(0) 0.0245(7)	0.40
0130	0.2586(2)	-0.00030(8) 0.63084(10)	-0.9124(2) 0.44014(17)	0.0243(7)	0.40
014a 014b	0.2380(2) 0.7502(2)	0.03984(10) 0.14100(10)	0.44014(17) 0.42057(17)	0.0233(0)	0
0140	0.7392(2)	-0.14199(10)	-0.43937(17)	0.0241(0)	0 15
015h	0.3416(2) 1.0107(2)	0.41004(11) 0.00405(12)	1.0177(2)	0.0292(7)	0.15
	1.0197(3)	0.09403(13)	-1.00000(18)	0.0555(8)	0.01
п9а1 110-2	0.205(2)	0.0824(19)	0.165(4)	0.049(3)	
H982	0.379(3)	0.0480(15)	0.191(4)	0.049(3)	
ПУDI Ц10-2	0.709(3)	-0.188(3)	-0.199(4)	0.049(3)	
H10a2	0.021(4)	0.7119(18)	0.052(4)	0.049(3)	
	1.018(3)	-0.2303(19)	-0.042(3)	0.049(3)	
H1002	1.122(4)	-0.208/(1/)	-0.044(3)	0.049(3)	
H11a1	0.721(2)	0.804(2)	0.320(4)	0.049(3)	
п11a2 1111ь1	0.024(4)	0.8440(15)	0.292(4)	0.049(3)	
	1.232(2)	-0.5130(19)	-0.302(4)	0.049(3)	
n1102	1.111(3)	-0.3442(14)	-0.301(4)	0.049(3)	

Table 4. Refined atom coordinates, equivalent isotropic displacement parameters ($Å^2$) and bond valence sums (*BVSs*, in valence units) for fluor-rewitzerite.

Table 4. Continued.

Atom	x	у	z	Ueq
H12a1	0.467(3)	0.7468(17)	0.439(4)	0.049(3)
H12a2	0.359(4)	0.7142(14)	0.403(4)	0.049(3)
H12b1	0.856(4)	-0.2208(17)	-0.415(4)	0.049(3)
H13a1	0.627(4)	0.4714(18)	0.938(3)	0.049(3)
H13a2	0.737(3)	0.5089(18)	0.944(4)	0.049(3)
H13b1	1.223(4)	0.0022(18)	-0.960(3)	0.049(3)
H13b2	1.097(3)	0.019(2)	-0.936(3)	0.049(3)
H14a1	0.310(4)	0.6195(19)	0.483(3)	0.049(3)
H14a2	0.245(4)	0.6179(17)	0.386(3)	0.049(3)
H14b1	0.751(4)	-0.1178(16)	-0.385(2)	0.049(3)
H14b2	0.796(4)	-0.1202(17)	-0.492(3)	0.049(3)
H15a1	0.468(3)	0.405(3)	0.987(4)	0.049(3)
H15b1	0.948(3)	0.089(2)	-0.976(3)	0.049(3)
H15b2	1.018(3)	0.078(2)	-1.070(2)	0.049(3)
Hala	0.685(5)	0.877(2)	0.014(4)	0.049(3)



Figure 6. The (001) section through the fluor-rewitzerite structure at z = 0. Dashed red lines on the left-hand side of the figure show H bonds, and solid blue lines on the right-hand side show bonding associated with the A1 and A2 sites. Arrows show [110] rows of regularly spaced H₂O groups.

Table 5. Refined site scattering for fluor-rewitzerite.

Site	Scattering used in refinement*	Scattering (electrons)
M1a	0.940(7) Mn + 0.060 Mg	24.2
<i>M</i> 1b	0.946(8) Mn + 0.054 Mg	24.3
M2a	0.593(12) Al+0.407 Ti	16.7
M2b	0.476(12) Al + 0.524 Ti	17.7
M3a	0.753(15) Al + 0.247 Ti	15.2
<i>M</i> 3b	0.685(15) Al + 0.315 Ti	15.8
A1	$0.130(7) \text{ K} + 0.770(7) \text{ O} + 0.1 \square$	8.6
A2	$0.887(6) \text{ K} + 0.013 \text{ O} + 0.1 \square$	17.0

* Mn scattering curve used for Mn plus minor Fe; Ti scattering curve used for Ti plus minor Fe.

X–O3–O7 of an M3-centred octahedron and forms H bonds with all three anions forming the face as shown in Fig. 6.

The major difference between the crystal structure of fluor-rewitzerite and that of the orthorhombic paulkerritegroup mineral benyacarite (Demartin et al., 1993) is an ordering of K and H₂O at the A1 and A2 sites, whereas they are disordered at a single A site in benyacarite. It would be expected that the ordering of the positive K⁺ cation at a specific A site would be associated with a localised negative charge at the associated structural unit site. For the ideal end-member fluor-rewitzerite with a structural unit formula of [Mn₂Al₂Ti(PO₄)₄(OF)(H₂O)₁₀]¹⁻, the local M2–X–M3ordered groupings in the structural unit will be predominantly Al–F–Al and Al–O–Ti. The first is formally chargebalanced for octahedrally coordinated Al, but the latter will give rise to an undersaturated bridging O^{2–} anion. It would

	U^{11}	<i>U</i> ²²	U ³³	U^{12}	<i>U</i> ¹³	U ²³
M1a	0.0214(3)	0.0096(2)	0.0212(3)	0.00151(15)	0.00041(17)	0.00117(14)
<i>M</i> 1b	0.0220(3)	0.0100(2)	0.0232(3)	-0.00056(15)	0.00029(17)	0.00246(14)
M2a	0.0190(5)	0.0090(4)	0.0295(5)	-0.00007(18)	-0.0012(3)	-0.00008(18)
M2b	0.0212(5)	0.0090(4)	0.0283(4)	-0.00042(17)	0.0014(3)	-0.00075(17)
M3a	0.0255(8)	0.0110(6)	0.0200(6)	0.0000(3)	0.0024(4)	0.0001(3)
<i>M</i> 3b	0.0242(7)	0.0104(5)	0.0185(6)	-0.0003(3)	-0.0011(4)	-0.0001(3)
P1a	0.0211(4)	0.0092(3)	0.0181(3)	-0.0001(2)	0.0007(2)	0.0016(2)
P1b	0.0213(4)	0.0094(3)	0.0193(3)	-0.0002(2)	-0.0001(2)	0.0012(2)
P2a	0.0208(4)	0.0089(3)	0.0192(3)	-0.0001(2)	0.0000(2)	0.0010(2)
P2b	0.0214(4)	0.0090(3)	0.0187(3)	0.0001(2)	0.0004(2)	0.0019(2)
A1	0.0476(15)	0.0153(10)	0.0314(12)	0.0043(7)	-0.0027(8)	-0.0014(6)
A2	0.0279(4)	0.0139(3)	0.0238(4)	-0.0003(2)	0.0026(2)	0.0018(2)

Table 6. Anisotropic displacement parameters $(Å^2)$ for metal atom sites in fluor-rewitzerite.

 Table 7. Hydrogen bonding in fluor-rewitzerite.

$D-\mathrm{H}\cdots A$	<i>D</i> –H (Å)	$\operatorname{H}\!\cdots\!A$ (Å)	$D \cdots A$ (Å)	∠ <i>D</i> H <i>A</i> (°)
O9a−H9a1···O5b	0.86(3)	1.83(3)	2.684(3)	170(4)
O9a−H9a2···O6a	0.84(3)	1.99(3)	2.835(3)	177(4)
O9b–H9b1···O5a	0.85(3)	1.89(3)	2.694(3)	157(5)
O10a-H10a2····O14b	0.84(4)	2.03(4)	2.807(3)	153(4)
O10b-H10b1···O12b	0.84(3)	2.02(4)	2.787(3)	153(4)
O10b-H10b2···O14a	0.86(4)	2.01(4)	2.758(3)	145(4)
O11a–H11a1…O1a	0.82(3)	2.01(3)	2.773(3)	156(4)
O11a–H11a2····O2b	0.83(3)	2.05(3)	2.854(3)	161(4)
O11a–H11a2····O4a	0.83(3)	2.50(4)	3.048(3)	125(3)
O11b-H11b1···O1b	0.89(3)	1.82(3)	2.709(3)	173(5)
O11b–H11b2···O2a	0.83(3)	2.01(3)	2.832(3)	170(5)
O12a−H12a1···O10a	0.83(4)	2.06(4)	2.809(3)	150(4)
O12a−H12a2···O14a	0.86(3)	1.91(3)	2.740(3)	162(4)
O12b-H12b1···O14b	0.82(4)	1.93(4)	2.748(3)	174(4)
O13a−H13a1····O15a	0.85(4)	1.82(4)	2.661(3)	170(4)
O13a–H13a2···X2	0.84(4)	2.27(4)	3.025(3)	149(4)
O13a–H13a2····O3a	0.84(4)	2.16(4)	2.831(3)	137(4)
O13a–H13a2···O7b	0.84(4)	2.45(4)	2.993(3)	123(3)
O13b-H13b1···X1	0.86(4)	2.23(4)	3.079(4)	173(4)
O13b-H13b1···O3b	0.86(4)	2.46(4)	2.918(3)	114(3)
O13b–H13b1···O7a	0.86(4)	2.43(4)	2.944(3)	119(3)
O13b-H13b2···O15b	0.85(4)	1.93(4)	2.697(3)	150(4)
O14a–H14a1····O3b	0.85(4)	1.92(4)	2.766(3)	174(4)
O14a–H14a2···O8b	0.81(3)	1.96(3)	2.775(3)	177(4)
O14b–H14b1···O8a	0.84(3)	1.94(3)	2.771(3)	172(4)
O14b-H14b2···O3a	0.87(3)	1.96(4)	2.812(3)	165(3)
O15b–H15b1····O14a	0.84(3)	2.44(3)	3.155(4)	144(4)
O15b–H15b2···O2a	0.84(3)	2.12(3)	2.918(3)	159(4)
A1–Ha1a…O7a	0.84(5)	1.97(5)	2.806(3)	171(5)

be expected that the K^+ would be associated with this configuration, and this appears to be borne out by the results in Tables 5 and 8. The K at the A2 site is coordinated to X2, which is shared between M2b containing dominant Ti and M3b containing dominant A1, whereas H₂O at the X1 site is associated with M2a and M3a, both containing dominant Al. This ordering should be manifested in X2 being predominantly O^{2-} and X1 being predominantly F⁻, but the *BVS* values in Table 4 do not reflect this. However, the *BVS* values for X1 and X2 are calculated from the mean positions of cations at the M2 and M3 sites, whereas these sites contain different cations, Al³⁺, Fe³⁺ and Ti⁴⁺, which may each be

M1a–O1b	2.084(2)	M1b–O1a	2.086(2)
–O5a	2.088(2)	–O5b	2.081(2)
O9a	2.178(2)	–O9b	2.153(2)
-O10a	2.221(2)	-O10b	2.238(2)
O11a	2.178(2)	O11b	2.159(2)
-O12a	2.208(2)	-O12b	2.257(2)
Avg.	2.159	Avg.	2.163
M2a–X1	1.777(2)	M2b-X2	1.774(2)
–O2b	1.966(2)	–O2a	1.978(2)
–O4a	1.906(2)	–O4b	1.904(2)
-O6a	1.929(2)	-O6b	1.934(2)
-O8b	1.915(2)	-O8a	1.911(2)
-O13a	2.050(3)	-O13b	2.079(3)
Avg.	1.924	Avg.	1.930
M3a–X1 x2	1.873(2)	M3b–X2 x2	1.897(2)
-O3b x2	1.913(2)	-O3a x2	1.909(2)
–O7a x2	1.915(2)	–O7b x2	1.892(2)
Avg.	1.900	Avg.	1.899
P1a–O1a	1.523(2)	P1b–O1b	1.529(2)
–O2a	1.545(2)	–O2b	1.540(2)
–O3a	1.542(2)	–O3b	1.530(2)
-O4a	1.539(2)	–O4b	1.540(2)
Avg.	1.537	Avg.	1.535
P2a–O5a	1.525(2)	P2b–O5b	1.528(2)
-O6a	1.534(2)	-O6b	1.535(2)
–O7a	1.533(2)	–O7b	1.537(2)
–O8a	1.553(2)	-O8b	1.541(2)
Avg.	1.536	Avg.	1.535
A1-X1	3.132(2)	A2–X2	3.070(2)
–O4a	2.815(3)	–O4b	2.858(2)
-O7a	2.801(3)	-O7b	2.771(2)
O9b	3.293(3)	O9a	3.200(2)
-O10a	2.673(3)	-O10b	2.720(2)
-O11a	3.350(3)	-O11b	3.251(3)
-O12b	2.835(3)	-O12a	2.836(2)
O15b	3.056(3)	-O15a	2.754(3)
Avg.	2.994	Avg.	2.933

 Table 8. Polyhedral bond lengths [Å] for fluor-rewitzerite.

at different positions in the octahedra. The anisotropic displacement parameters (ADPs) for the M2 sites in Table 6 show elevated values for U^{33} , corresponding to root-meansquare displacements of 0.17 Å. Displacements of individual cations by this magnitude will have large effects on the BVS values for the X sites. We tried refining the O/F contents at these sites, but these anions are isoelectronic and the refinements were inconclusive. We also tried refinements in the non-centrosymmetric space group Pc, but the refinements were unstable (oscillating) and gave non-positive ADPs and large ranges for the P–O distances.

first three members to be characterised, paulkerrite (Peacor et al., 1984), mantienneite (Fransolet et al., 1984) and benyacarite (Demartin et al., 1993, 1997), were all described as having orthorhombic symmetry, space group Pbca, and this appeared to be confirmed from a single-crystal refinement for benyacarite. The new members pleysteinite and hochleitnerite (Grey et al., 2023c, d) were reported to be isostructural with benyacarite, based on laboratory-based single-crystal diffraction data. A single-crystal study on type paulkerrite using microfocus synchrotron diffraction data, however, revealed a small monoclinic distortion of the published unit cell ($\beta = 90.33(3)^{\circ}$) with the crystal structure conforming to $P2_1/c$, a maximum non-isomorphous subgroup of *Pbca* (Grey et al., 2023b). Peacor et al. (1984) had already noted the poor diffracting quality of paulkerrite crystals and the "possibility of submicroscopic domains having a structure differing from the average". We subsequently collected microfocus synchrotron diffraction data on crystals of pleysteinite and hochleitnerite and confirmed that these minerals are also monoclinic, $P2_1/c$ (Rewitzer et al., 2024). The combination of a very small deviation of β from 90°, ubiquitous twinning and large mosaicity makes it difficult to determine the correct symmetry using conventional sealed-tube laboratory diffractometer data. Recent microfocus synchrotron data collected on the new paulkerrite-group minerals, rewitzerite (Grey et al., 2023a), hydroxylbenyacarite (Hochleitner et al., 2024) and macraeite (Grey et al., 2024), have confirmed the monoclinic character of these minerals. It is likely that the crystal structure of benyacarite is also of monoclinic symmetry. An indication of this comes from the very large displacement parameters reported for O15 (Demartin et al., 1993), with $U \sim 0.05 \text{ Å}^2$ corresponding to a mean-square displacement of 0.22 Å. In going from the average *Pbca* structure to $P2_1/c$, the largest displacements occur for O15 in published monoclinic structures (Rewitzer et al., 2024). This is shown for fluor-rewitzerite in Table 8, where A1-O15b and A2–O15a distances differ by 0.30 Å, whereas all the other pairs of A–O distances differ by less than 0.1 Å. The large U values for O15 from the Pbca refinement of benyacarite thus suggest that the correct symmetry is monoclinic, at least in small domains.

There are currently nine published paulkerrite-group minerals, which are distinguished compositionally in Fig. 4. The

Some trends in the crystal chemistry of the seven monoclinic paulkerrite-group minerals are given in Table 9. The magnitude of the monoclinic distortion correlates positively with the difference in scattering between the A1 and A2 sites, that is, with the extent of ordering of K and H₂O at the A sites. This is illustrated in Fig. 7 ($R^2 = 0.70$). There is a correlation between the monoclinicity and the extent of twinning. The minerals with the greatest monoclinicity are untwinned or have a dominant twin individual, and decreasing monoclinicity is associated with a shift towards equal twin volumes. Also given in Table 9 is information on the reflections (*hk*0), h = 2n + 1, and (0*kl*), k = 2n + 1. These

	Pleysteinite	Fluor- rewitzerite	Paulkerrite	Rewitzerite	Macraeite	Hochleitnerite	Hydroxyl- benyacarite
$A1, A2, M1, (M2_2M3), X_2$	H ₂ O, K, Mn, (Al ₃), F ₂	H ₂ O, K, Mn, (Al ₂ Ti), OF	H ₂ O, K, Mg, (Fe ₂ Ti), OF	K, H ₂ O, Mn, (Al ₂ Ti), O(OH)	H ₂ O, K, Mn, (Fe ₂ Ti), O(OH)	H ₂ O, K, Mn, (Ti ₂ Fe), O ₂	H ₂ O, H ₂ O, Mn, Ti ₂ Fe, (O(OH)
K ₂ O (wt %)	5.20	5.78	5.15	3.93	4.14	4.64	2.25
TiO ₂ (wt %)	5.47	7.64	10.85	9.18	9.85	12.95	14.57
Al ₂ O ₃ (wt %)	12.14	9.27	1.29	8.37	1.65	0.64	0.08
Fe ₂ O ₃ (wt %)	1.72	3.54	12.66	7.44	11.80	9.78	9.98
F (wt %)	2.41	1.88	0.99	0.87	1.04	0.44	0.32
a (Å) b (Å) c (Å) β (°)	10.440(5) 20.588(5) 12.234(2) 90.38(1)°	10.407(1) 20.514(2) 12.193(1) 90.49(2)	10.569(2) 20.590(4) 12.413(2) 90.33(3)°	10.442(2) 20.445(2) 12.269(1) 90.17(3)°	10.562(2) 20.725(4) 12.416(2) 90.09(3)°	10.547(2) 20.577(4) 12.373(2) 90.09(3)°	10.5467(3) 20.7222(5) 12.5031(3) 90.068(2)
Twin volumes	0.91, 0.09	0.80, 0.20	No twinning	0.62, 0.38	0.63, 0.37	0.55, 0.45	0.51, 0.49
Site scattering A1, A2 (electrons)	18.2, 7.7	17.1, 8.9	15.6, 8.7	15.7, 8.3	15.2, 9.9	12.8, 8.5	9.8, 7.7
*Observed, all, R_p hk0, h = 2n + 1	85, 173, 0.12	112, 152, 0.10	40, 168, 0.19	101, 130, 0.08	87, 119, 0.13	66, 146, 0.16	55, 143, 0.28
0kl, k = 2n + 1	104, 185, 0.12	92, 106, 0.14	37, 188, 0.17	100, 121, 0.11	94, 136, 0.13	94, 160, 0.15	75, 200, 0.28

Table 9. Crystal-chemical trends in monoclinic paulkerrite-group minerals, general formula $A1A2M_1M_2M_2M_3(PO_4)_4X_2(H_2O)_{10} \cdot 4H_2O$.

* Reflections forbidden in Pbca.



Figure 7. Plot of *A*-site scattering difference vs. β -90 (°) for monoclinic paulkerrite-group minerals.

reflections are systematically absent in the diffraction patterns of the orthorhombic *Pbca* minerals due to the *a* and *b* glide planes. Their presence, as well as their fitting in the refinement, is the most direct evidence for the space group $P2_1/c$. The results for paulkerrite appear to be anomalous, with much lower numbers of these reflections, but this is because the reflection data are much weaker, with only 40% of reflections having $I > 3\sigma(I)$ (Grey et al., 2023b) compared with ~ 90% for the other minerals. Aside from paulkerrite, there is a general trend from left to right (direction of decreasing monoclinicity and K/H₂O ordering) of smaller numbers of observed "forbidden" reflections and increasing partial *R* factors (R_p) in the fitting of these reflections. Hydroxylbenyacarite, which has only about half the K content of the other minerals, is the most extreme in terms of the smallest monoclinicity and *A*-site scattering difference, and it has the poorest fit to the forbidden reflection intensities. Reconstructed precession images of the diffraction data for this mineral show that the reflections (hk0), h = 2n + 1, and (0kl), k = 2n + 1, are diffuse with streaking parallel to b^* (Hochleitner et al., 2024). The lengths of the diffuse streaks correspond to correlation lengths of only 1 to 2 nm, so the *A*site ordering is restricted to domains on the scale of the unit cell. The results in Table 9 are consistent with a progressive decease in the size of the domains in which ordering of K and H₂O at the *A* sites occurs, in the minerals from left to right in the table.

The general crystallographic trends shown in Table 9 correlate strongly with the mineral compositions. In particular, the decrease in *A* site scattering and associated smaller monoclinicity from left to right in Table 9 are associated with decreases in K, Al and F and with increases in Ti content. The variation in Fe³⁺ is less clear, but overall it increases from left to right, with the clear exception of paulkerrite. The elements Al, Fe³⁺, Ti and F are all associated with the octahedra at the $M2_2M_3$ sites, suggesting that it is the distribution of compositions at these sites that determines the extent of monoclinic ordering. Figure 4 shows a distinct difference in the compositions of minerals that have Al > Fe at the $M2_2M_3$ sites compared with those that have Fe > Al. The former all have $M2_2M_3$ compositions in the range of M_2^{3+} Ti to M_3^{3+} (M^{3+} -rich), whereas the latter have compositions in the range of M_2^{3+} Ti to M^{3+} Ti₂ (Ti-rich). Minerals in the Alrich Al₃ and Al₂Ti phase fields (pleysteinite, fluor-rewitzerite and rewitzerite) have the highest degree of monoclinic ordering, whereas minerals in the Ti-rich Ti₂Fe phase field have the lowest degree of ordering (hydroxylbenyacarite, hochleitnerite and benyacarite). The mineral mantienneite remains unclassified because the crystals are not of high-enough quality for a single-crystal structure determination. At present there are no known paulkerrite-group minerals with $M2_2M_3$ compositions in the Fe₃ or Ti₂Al phase fields. Eventual discovery and characterisation of minerals in these two phase fields will help to further develop an understanding of the factors that determine the monoclinic ordering.

Data availability. Crystallographic data for fluor-rewitzerite are available in the Supplement.

Supplement. The supplement related to this article is available online at: https://doi.org/10.5194/ejm-36-541-2024-supplement.

Author contributions. IEG oversaw the research and wrote the paper; RH conducted initial characterisation studies on the specimen and identified it as a potential new species; WGM assisted in the diffraction data analysis; ARK measured the optical properties, Raman spectrum, powder XRD and crystal morphology; SB collected and processed the single-crystal diffraction data; and NCW and CMM developed the cryo-EMP analysis procedure and associated software.

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